
Magnetic mineralogy of Variscan granites from northern Portugal: an approach to their petrogenesis and metallogenic potential

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| A B S T R A C T |

Northern Portugal is characterized by the occurrence of numerous W hydrothermal deposits spatially associated with granites. The primary goal of this work is to establish a relationship between the magnetic behavior of the granites and the redox conditions during magma genesis, as this can influence the occurrence of mineralizations, namely of W (Mo). To this end, the magnetic mineralogy of the granites of the Lamas de Olo Pluton, a post-tectonic pluton in northern Portugal, with associated W (Mo) occurrences was characterized and compared with the magnetic mineralogy of other post-tectonic Variscan plutons. This pluton is composed of different biotite granites: Lamas de Olo, Alto dos Cabeços and Barragem. To better characterize its magnetic behavior, different analytical techniques that complement previous magnetic susceptibility studies were performed. The magnetic mineralogy of Lamas de Olo Pluton was then compared with other post-tectonic Variscan plutons such as the Vila Pouca de Aguiar, Peneda-Gerês and Lavadores-Madalena plutons. The presence of magnetite in some of these granites is important because it points to melt-oxidized conditions not commonly found in Iberian Variscan granites. Our study shows that granite areas where magnetite and/or magnetite/ilmenite coexist are important targets for W (Mo) mineralizations. The results indicate that a few plutons have granites with a complex redox history which leads to the formation of magnetite and ilmenite.

KEYWORDS

Magnetic mineralogy. Variscan granites. Post-D₃ biotite granites. Central Iberian Zone. W-Mo mineralizations.

INTRODUCTION

The northern area of Portugal is characterized by large volumes of granitic intrusions related to Variscan orogeny (321-290Ma). Synorogenic granites represent the plutonic magmatism and, based on their geological, petrographic and geochemical characteristics, are divided into two main groups (e.g. Chappell and White, 1974; Dias *et al.*, 2010; Ferreira *et al.*, 1987; Noronha *et al.*, 2006). The first group consists of two-

mica peraluminous granites, which are dominantly syntectonic (syn-D₃: 321-312Ma) and considered S-type granites, resulting from the crystallization of wet peraluminous magmas originated at a mesocrustal level. The second group consists of biotite granites considered to have been generated at deep crustal levels and corresponding to dry magmas. The age of their formation is either syn-D₃ (321-312Ma), late-D₃ (312-305Ma), late- to post-D₃ (ca. 300Ma), or post-D₃ (299-290Ma) (Dias *et al.*, 2010).

These two types of granites have distinctive whole-rock oxygen isotope ($\delta^{18}\text{O}$) values. The highest $\delta^{18}\text{O}$ values indicate a crustal origin, and the lowest values suggest a mantle contribution and/or source region dominated by mafic and meta-igneous rocks (Ellwood and Wenner, 1981; Ishihara, 1977). The $\delta^{18}\text{O}$ mean values range between 10.65‰ and 12.90‰ for the two-mica granites and 9.75‰ and 12.84‰ for the biotite granites (Antunes *et al.*, 2008; Cruz *et al.*, 2016; Sant’Ovaia *et al.*, 2013a, b; Teixeira *et al.*, 2012).

The magnetic susceptibility (K_m) of granites is an important characteristic and it is mainly controlled by the presence of certain oxide minerals like magnetite and/or ilmenite as well as ferromagnesian silicates such as biotite.

The abundance of magnetite or ilmenite can be explained by different redox conditions in the magma chamber and different magma sources. The presence of magnetite or ilmenite as accessory minerals represents oxidized- or magnetite-type granites and reduced- or ilmenite-type granites, respectively (Ishihara, 1977). Magnetite granites are considered to have been generated at great depth (upper mantle and/or lower crust), whereas the ilmenite-series are considered to have originated at a shallower level (middle to lower continental crust) where small amounts of crustal carbon are present (Ellwood and Wenner, 1981; Ishihara, 1977; Sheppard, 1977). Relative abundances of magnetic minerals in granites can be measured in terms of magnetic susceptibility, with the magnetite-series having $K_m > 3.0 \cdot 10^{-3}\text{SI}$ and ilmenite-series, in general, with K_m around 10^{-6}SI (or μSI) (e.g. Ishihara, 1977; Takagi and Tsukimura, 1997).

In the Iberian Variscan belt, specifically in the Central Iberian Zone (CIZ), in Spain and Portugal, the magnetic behavior of numerous granites has been analyzed. The Spanish granites are mostly ilmenite-type and have been extensively analyzed by several authors, like Aranguren *et al.* (1996); Olivia-Urcia *et al.* (2012); Porquet *et al.* (2017); Román-Berdiel *et al.* (1995). These authors found K_m values between $22\mu\text{SI}$ and $467\mu\text{SI}$. Recent studies carried out by Villaseca *et al.* (2017) proved the existence of both ilmenite- and magnetite-type granites in the Spanish Central System. The magnetic susceptibility of these granites yielded K_m ca. $15\mu\text{SI}$ to $180\mu\text{SI}$ for S- and I-type granites (ilmenite-type granites) and values between $500\mu\text{SI}$ and $1,400\mu\text{SI}$ for the leucogranites I-type granites (magnetite-type granites). Magnetic susceptibility studies in several Portuguese granites (Sant’Ovaia *et al.*, 2014) yielded K_m mean values between $48\mu\text{SI}$ and $84\mu\text{SI}$ for two-mica granites corresponding to ilmenite-type granites and between $72\mu\text{SI}$ and $11,676\mu\text{SI}$ for biotite granites corresponding to ilmenite-type and/or magnetite-type granites.

Several metallogenic events are identified on the granite bodies: i) Sn-Li pegmatites are essentially associated to syn-D₃ with two-mica granites; ii) W (Sn) and W quartz-vein deposits occur in late- and late- to post- D₃ biotite granites; and iii) W (Mo) deposits are related to post-D₃ biotite sub-alkaline granites (e.g. Mateus and Noronha, 2010; Noronha, 2017; Thadeu, 1965).

The magnetic susceptibility (K_m) and $\delta^{18}\text{O}$ data allow to establish a relationship between the granite type (ilmenite- or magnetite-type) and the associated mineralizations (Kumar, 2010; Takagi and Tsukimura, 1997). The occurrence of ore deposits associated with ilmenite-type and/or magnetite-type granites has been largely described, for example, in Indonesia (Maulana *et al.*, 2013) and, in southern Korea and southwestern Japan (Ishihara *et al.*, 1981) and, more recently, in Portugal (Cruz *et al.*, 2016; Sant’Ovaia *et al.*, 2014). In Sulawesi, Indonesia, the ilmenite-type granites outcrop in the southern zone of the island while the magnetite-type granites occur in the northern area. Both granites have associated ore mineralization. Cu-Au-Mo mineralizations are associated with magnetite-series granitic rocks, while Sn-W mineralizations related to reduced ilmenite-type granites have not yet been reported (Maulana *et al.*, 2013). Contrastingly, in southern Korea and southwestern Japan (Ishihara *et al.*, 1981), W-Mo occurrences were described in areas with both ilmenite- and magnetite-type granites with no clear separation in the distribution of the two types of granites. In the Portuguese sector of the Central Iberian Zone, the Sn deposits are associated with S- and ilmenite-type granites and the W(Sn) mineralizations occur in veins that cut the I (peraluminous)- and ilmenite-type granites, both with high $\delta^{18}\text{O}$ values. The W (Mo) occurrences are spatially related to I (peraluminous)- and magnetite-type granites (Cruz *et al.*, 2016; Sant’Ovaia *et al.*, 2012).

The study of magnetic mineralogy expands our knowledge of the relation between granites and their related mineralizations, since the redox conditions control the specific mineral occurrences. It is very important to understand the relationship between magnetic mineralogy and ore deposits considering the complexity of magma genesis.

The main goal of this work is to characterize the magnetic behavior of the Lamas de Olo Pluton (LOP) and compare it with the magnetic behavior of other post-tectonic Variscan biotite plutons (Peneda-Gerês Pluton (PGP), Lavadores-Madalena Pluton (LMP) and Vila Pouca de Aguiar Pluton (VPAP). We aim at establishing a relationship between the magnetic behavior of the granites and the occurrence of different kinds of mineralizations.

GENERAL FRAMEWORK

The Iberian Variscan belt

The Iberian Variscan belt was formed as a large curved segment of the European Variscan belt resulting from the Laurussia-Gondwana collision during Devonian to Carboniferous times (Franke, 1989; Kroner and Romer, 2013; Ribeiro *et al.*, 1990). After the collision, during the late Paleozoic, the Pangea amalgamation leads to continuous rock deformation and the C-shaped Cantabrian orocline was formed (e.g. Gutiérrez-Alonso *et al.*, 2012; Weil *et al.*, 2013). Other authors suggest a more complex scenario, with the existence of two oroclines, in S-shaped, the Cantabrian, and Central-Iberian (e.g. Pastor-Galán *et al.*, 2012; Shaw *et al.*, 2012).

The Iberian belt is subdivided into several geotectonic zones (Fig. 1): the Cantabrian Zone (CZ), the West Asturian Leonese Zone (WALZ), the CIZ, the Ossa Morena Zone

(OMZ) and the South Portuguese Zone (SPZ) (Julivert *et al.*, 1974). Farias *et al.* (1987) identified allochthonous and parautochthonous units in the CIZ, which led them to consider another zone, the Galicia Trás-os-Montes Zone (GTMZ).

Three main ductile deformation phases have been documented in northern Portugal, namely D₁, D₂ and D₃ (Noronha *et al.*, 1981). In the CIZ, two main phases of deformation can be recognized in the autochthonous terrane: D₁ (ca. 360–337Ma, e.g. Castiñeiras *et al.*, 2008; Dallmeyer *et al.*, 1997; Díez Fernández *et al.*, 2016; Martínez Catalán *et al.*, 2014; Pereira *et al.*, 2018), and D₃ (ca. 321–300Ma, e.g. Ferreira *et al.*, 1987; Pereira *et al.*, 2018). D₁ produced isoclinal folds with a subvertical axial plane with a NW-SE strike and a well-marked axial plane schistosity (S₁), and D₃ is characterized by wide folding of small amplitude on the subvertical axial plane and sub-horizontal NW-SE axis. Simultaneously with the D₃, vertical ductile shear zones were developed in lower structural levels. In the upper

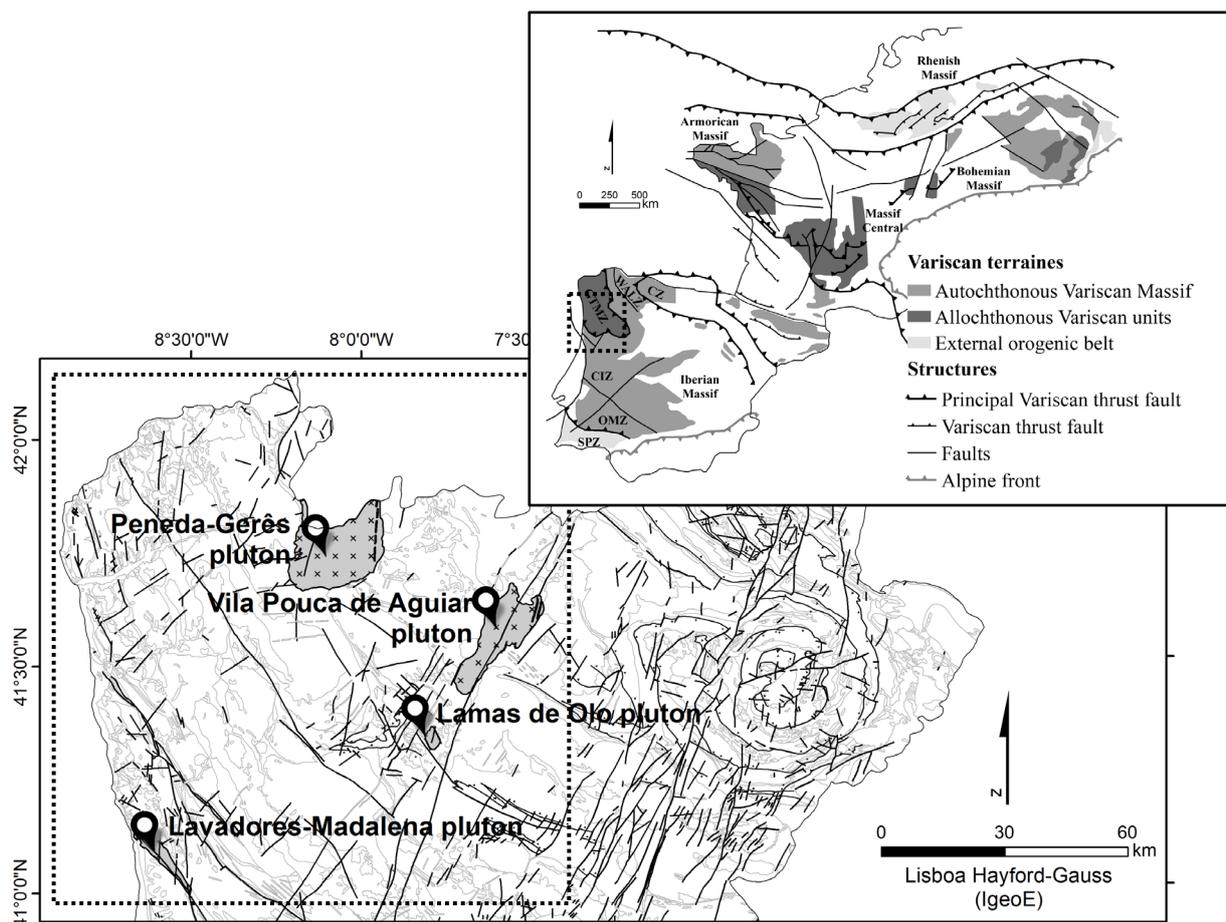


FIGURE 1. Iberian Variscan belt and location of the studied granitic bodies: Lamas de Olo Pluton, Vila Pouca de Aguiar Pluton, Peneda-Gerês Pluton and Lavadores-Madalena Pluton (modified from Ferreira *et al.*, 1987). CZ: Cantabrian Zone, WALZ: West Asturian Leonese Zone, GTMZ: Galicia Trás-os-Montes Zone, CIZ: Central Iberian Zone, OMZ: Ossa Morena Zone, SPZ: South Portuguese Zone (based on: Farias *et al.*, 1987; Julivert *et al.*, 1974).

structural levels, a brittle deformation occurred, leading to fracture-conjugate systems in a NNE-SSW and NNW-SSE direction (Marques *et al.*, 2002). D₂ (ca. 337-316Ma; e.g. Dallmeyer *et al.*, 1997; Díez Fernández *et al.*, 2016; Martínez Catalán *et al.*, 2014; Pereira *et al.*, 2018) is well represented in allochthonous terranes in the GTMZ.

Most of the granite intrusions are coeval with D₃ and the others are late- and post-D₃ (e.g. Dias *et al.*, 2010; Ferreira *et al.*, 1987). The development of different tectonic structures, namely shear zones, during D₃ and successive reactivation events controlled the granite emplacement and caused extensive hydrothermal activity throughout the entire crust, involving distinct fluid sources, some transporting significant amounts of ore mineral phases (Mateus and Noronha, 2010; Noronha *et al.*, 2013).

Plutons and associated mineralizations

The LOP is located in the northern part of the CIZ near the limit with the GTMZ (Figs. 1; 2). It is a small post-tectonic pluton with a rhombus shape, controlled by NNW-SSE offsets conjugated with the NNE-SSW fault system, which are parallel to the Verin-Régua-Penacova fault (VRPF). The LOP is composed of distinct outcropping granites (Fig. 2): i) Lamas de Olo (LO), the more representative, a medium- to coarse-grained porphyritic granite (biotite > muscovite); ii) Alto dos Cabeços (AC), a fine- to medium-grained porphyritic granite (biotite > muscovite) and iii) Barragem (BA), a fine- to medium-grained slightly porphyritic leucocratic granite (biotite = muscovite) (Fernandes *et al.*, 2013; Pereira, 1989). ²⁰⁷Pb/²³⁵U dating in monazite yield an age of 297.19±0.73Ma for the LO granite (Fernandes *et al.*, 2013). Field observations show that the contact between the LO and AC granites is generally diffuse, and that the BA granite crosscuts the LO and AC granites. Previous studies of anisotropy of magnetic susceptibility showed that LOP has two types of behavior granites: magnetite- and ilmenite-type (Cruz *et al.*, 2016). W-Mo (Sn) mineralizations occur, mostly in N 80° E sub-vertical quartz veins (Helal, 1992).

The LOP intrudes Lower Paleozoic formations (Armorican quartzite and schists of Upper Silurian to Middle Ordovician age), the Douro Group metasediments (Upper to Middle Cambrian) and two-mica syntectonic granites in the Vila Real Massif (Fig. 2). The Vila Real Massif is a syn-D₃ two-mica granite composite massif with associated Sn mineralizations in pegmatites, and W-Sn in hydrothermal quartz veins (Pereira, 1989). Monazite and zircon analysis yield a weighted average ²⁰⁷Pb/²³⁵U age of ca. 311±1Ma (Almeida *et al.*, 1998).

The other granites referred to in this study belong to Vila Pouca de Aguiar, Peneda-Gerês and Lavadores-Madalena plutons (Fig. 1), classified as post-tectonic biotite plutons.

Vila Pouca de Aguiar and Peneda-Gerês plutons are located in the GTMZ and Lavadores-Madalena Pluton on the NW border of the CIZ.

The Vila Pouca de Aguiar Pluton is a composite pluton with two main biotite granites: Vila Pouca de Aguiar Granite (VPAG) and Pedras Salgadas Granite (PSG). The two granites have a similar mineralogical composition but different texture, with VPAG displaying a coarse grain porphyritic texture and PSG a more equigranular texture (Martins and Noronha, 2006; Sant'Ovaia *et al.*, 2000). As regards their ages, PSG is younger than VPAG, with 297±14Ma and 298±9.1Ma, respectively (Martins *et al.*, 2009).

The Peneda-Gerês Pluton is also a composite pluton with three main granites: Gerês, Illa-Bouzagrado, and Carris. The Illa-Bouzagrado is a leucogranite (muscovite > biotite) and the other two are biotite granites (Cottard, 1979; Mendes and Dias, 2004). All the granites are contemporaneous and dated to ca. 290-296Ma (Dias *et al.*, 1998). Some deposits associated with the granites of this pluton, have been described, namely: Mo (Bouzagrado in Spain) and W (Mo-Sn-Bi) (Carris and Borrageiro in Portugal and Las Sombras in Spain) (Cottard, 1979; Cheilletz and Giuliani, 1982).

The Lavadores-Madalena Pluton, dated to 298±11Ma (Martins *et al.*, 2011), is composed of Lavadores and Madalena granites. The Lavadores granite is a porphyritic, coarse-grained biotite granite (Martins *et al.*, 2011) and the Madalena granite is a porphyritic medium- to coarse-grained biotite granite (Sant'Ovaia *et al.*, 2014). W (Mo) mineralization associated with granites of this pluton has been described (Thadeu, 1965; Teixeira and Perdigão, 1962).

MATERIAL AND METHODS

Representative samples of different Variscan granites were examined. The sampled granites were from the Lamas de Olo Pluton, Vila Pouca de Aguiar Pluton, Peneda-Gerês Pluton and Lavadores-Madalena Pluton (Table 1; Fig. 1). Special attention was given to the LOP (Fig. 2), where 67 sites were sampled for this study.

To attain the primary objective of this work, different analytical methods were used: optical microscopy, magnetic susceptibility measurements, isothermal remanent magnetization acquisition curves, and thermomagnetic experiments. Previous studies (Cruz *et al.*, 2016 for LOP; Martins and Noronha, 2006; Sant'Ovaia, 1993; Sant'Ovaia *et al.*, 2000 for Vila Pouca de Aguiar Pluton; Cottard, 1979; Mendes and Dias, 2004 for Peneda-Gerês Pluton; Martins *et al.*, 2011; Sant'Ovaia *et al.*, 2014 for Lavadores-

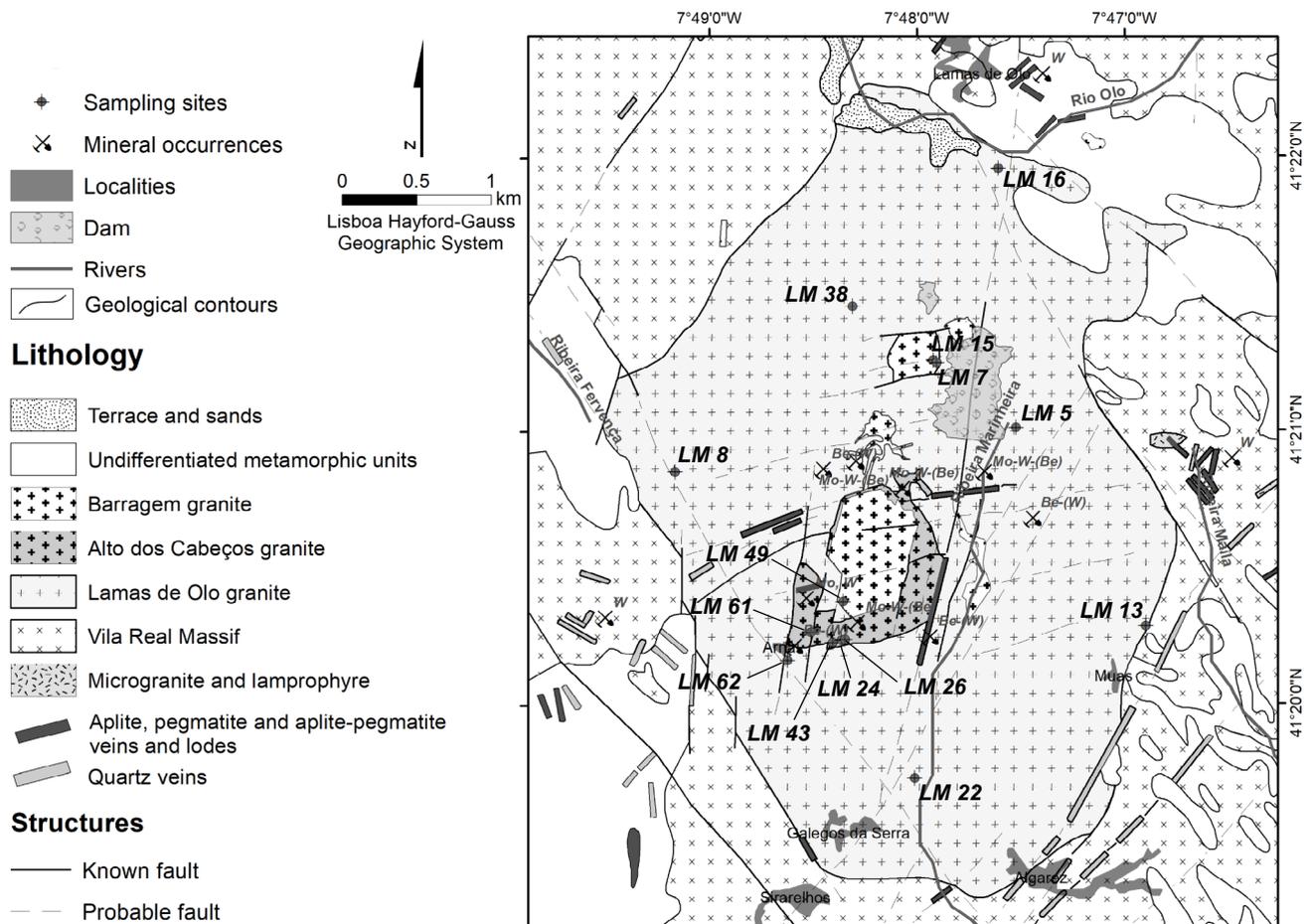


FIGURE 2. Simplified geological map from the Lamas de Olo Pluton (modified from Helal, 1992; Pereira et al., 1987; Pereira, 1989) and location of the samples (LM) used in the isothermal remanent magnetization acquisition curves, frequency-dependent susceptibility measurements and thermomagnetic experiments.

Madalena Pluton) were considered and their data compared with the new data (Table 1).

Microscopy

A Leica petrographic polarizing microscope with a digital camera was used for the reflected and transmitted light petrographic studies. The Raman spectra of different opaque minerals were obtained using a Raman LabRAM HORIBA Jobin Yvon Spex spectrometer interfaced with an Olympus microscope with 50x objective lens, diffraction gratings with 1,800 lines mm^{-1} and equipped with a 632.8nm emission line of a HeNe laser at a power of 20mW. The incident beam perpendicular to the plane of the sample was focused through the microscope lens, which also collected the Raman scattered radiation in back-scattering geometry. A highly sensitive Charge-Coupled Device (CCD) camera was used to collect the Raman spectra. Extended scans were performed on a spectral range from 50 to 1,200 cm^{-1} . The time of acquisition and the number of accumulations varied in order to obtain an optimized spectrum for each analyzed

mineral. Both equipment belong to the *Departamento de Geociências, Ambiente e Ordenamento do Território (DGAOT)* of the *Faculdade de Ciências da Universidade do Porto (FCUP)* and *Instituto de Ciências da Terra-Polo Porto (ICT-Porto)*.

Complementary studies to identify minor mineral amounts and characterize the magnetic mineralogy were carried out using a Scanning Electron Microscope (SEM) and Energy Dispersive Spectroscopy (EDS) X-ray microanalysis at the *Centro de Materiais da Universidade do Porto (CEMUP)*. The SEM/EDS studies were performed using a high-resolution JEOL JSM 6301F SEM coupled with an Oxford INCA Energy 350 system. Samples were carbon-coated using a JEOL JEE-4X Vacuum Evaporator.

A total of 13 thin sections from the LOP were used for petrographic studies (Table 1). For complementary studies, some representative samples were selected: 4 for Raman (3 from LO and 1 from BA) and 5 for SEM/EDS (3 from LO and 1 from BA). These samples were selected taking into

account the presence of magnetite, hematite, and ilmenite, which was previously identified under a microscope or by magnetic susceptibility studies. The results from other plutons were compiled from the literature (Table 1).

Magnetic susceptibility and isothermal remanent magnetization

The relationship between the magnetization induced in a material M and the external field H is defined as $M=KH$, where K is the magnetic susceptibility. The K reflects the whole-rock mineral composition, comprising the diamagnetic, paramagnetic and ferromagnetic minerals (Rochette, 1987). K measurements were performed using a KLY-4S Kappabridge susceptometer Agico model (Czech Republic) from the DGAOT-FCUP and ICT-Porto. Measurements were undertaken in a 300A/m field at room temperature. For each site, the ANISOFT 4.2 program package (Chadima and Jelinek, 2009) enabled us to calculate the mean susceptibility K_m . The magnetite is an important ferromagnetic mineral (*s.l.*), easily detected because increases the K_m values to $>10^{-3}$ SI (Bouchez, 1997, 2000; Tarling and Hrouda, 1993).

The magnetic susceptibility data were mostly compiled from the literature, and only a few samples from the Peneda-Gerês Pluton were measured in this study (Table 1).

For a detailed magnetic mineralogy study, the acquisition of Isothermal Remanent Magnetization (IRM) curves were performed. The IRM acquisition curves are important to estimate the characteristic coercivity of the ferromagnetic minerals (Butler, 1992). The IRM refers to the remanence acquired by a sample exposed to a direct magnetic field

(H), at ambient temperature. These measurements were performed in the *Laboratório de Paleomagnetismo* at the *Instituto Dom Luís, Lisboa*. Beforehand, standard samples were demagnetized through an alternating field cleaning at *ca.* 100mT, with an LDA-3A (Agico) demagnetizer. Then, the magnetic field (H) was imparted using an Impulse Magnetizer IM-10-30 (ASC Scientific) and the resulting remanence was measured using a spinner magnetometer JR-6A (Agico). Samples were magnetized in the same direction with increasing magnetic fields from *ca.* 3.51mT up to 1.15T in different steps (*ca.* 40 steps) and the induced magnetization acquired by the specimens was measured after each step of induction. Data were analyzed using a Cumulative Log-Gaussian (CLG) function (Robertson and France, 1994) with the software developed by Kruijver *et al.* (2001). For the IRM studies, 4 granite samples from the LOP were selected according to their magnetic behavior: 1 from the LO-magnetite-type granite, 1 from the LO-ilmenite-type granite, 1 from the BA granite and 1 from the AC granite (Table 1).

Additionally, measurements of frequency-dependent susceptibility were done with a Bartington MS2 System in the *Laboratório de Paleomagnetismo* at the *Instituto Dom Luís, Lisboa*. The low-field magnetic susceptibility was measured at two applied field frequencies (0.46kHz and 46kHz). The percentage of frequency-dependent susceptibility (KfD%) (Dearing *et al.*, 1996) was obtained in 10 samples from the LOP granites (Table 1).

Thermomagnetic experiments

The thermomagnetic experiments allow the identification of ferromagnetic minerals (*s.l.*) based on

TABLE 1. Methodologies applied to studied granites. LOP: Lamas de Olo Pluton; VPAP: Vila Pouca de Aguiar Pluton; PGP: Peneda-Gerês Pluton; LMP: Lavadores-Madalena Pluton; ✓: data obtained in this study; x: no data

Pluton	Granite	Methodologies				
		Optical microscopy studies	Magnetic susceptibility studies	Thermomagnetic curves	Isothermal remanent magnetization	Frequency-dependent susceptibility
LOP	Lamas de Olo	✓ (7 samples)		✓ (7 samples)	✓ (2 samples)	✓ (4 samples)
	Alto dos Cabeços	✓ (4 samples)	Cruz <i>et al.</i> (2016)	✓ (4 samples)	✓ (1 sample)	✓ (3 samples)
	Barragem	✓ (2 samples)		✓ (4 samples)	✓ (1 sample)	✓ (3 samples)
VPAP	Pedras Salgadas	Sant'Ovaia (1993) Sant'Ovaia <i>et al.</i> (2000)	Sant'Ovaia (1993) Sant'Ovaia <i>et al.</i> (2000)	✓ (1 sample)	Sant'Ovaia (1993)	x
	Vila Pouca Aguiar	Martins and Noronha (2006)		x		x
PGP	Gerês		✓ (7 samples)	✓ (1 sample)	x	x
	Illa	Mendes and Dias (2004) Cottard (1979)	✓ (3 samples)	x	x	x
	Carris		✓ (3 samples)	x	x	x
LMP	Madalena	Martins <i>et al.</i> (2011) Sant'Ovaia <i>et al.</i> (2014)	Sant'Ovaia <i>et al.</i> (2014)	✓ (1 sample)	Martins <i>et al.</i> (2011) Sant'Ovaia <i>et al.</i> (2014)	x
	Lavadores			✓ (1 sample)		x

their Curie/Néel temperature. The Curie temperature (T_C) is the temperature above which a ferromagnetic material (*s.s.*) becomes paramagnetic, and the Néel temperature (T_N) is analogous to the T_C but for antiferromagnetic materials.

The Curie/Néel temperature of ferromagnetic minerals (*s.l.*) can be determined in low-field thermomagnetic experiments where the magnetic susceptibility of a sample is monitored while temperature is increased or decreased, and it is defined as the point of major decrease in magnetic susceptibility of the sample during the heating cycle (Butler, 1992). The temperature of this drop in magnetic susceptibility corresponds to the temperature at which a magnetic mineral loses its spontaneous magnetization. In contrast, if magnetic susceptibility, with values much lower than ferromagnetic values, decreases regularly with increasing temperature, the sample has the typical behavior of paramagnetic minerals (Dunlop and Özdemir, 1997).

The Curie/Néel temperature is typical for a particular mineral and is therefore commonly used to identify its composition. For example, the Curie/Néel temperature is 580°C for magnetite, 680°C for hematite, 320°C for pyrrhotite and 120°C for goethite. However, the rocks can have other iron minerals with titanium in their composition, corresponding to the titanomagnetite and titanohematite series. It should be noted that, in the titanomagnetite series, the Curie point decreases as the Ti content increases; therefore, the T_C can range between 580°C (T_C of pure magnetite) and -150°C (T_C of pure ulvospinel) (Dunlop and Özdemir, 1997).

The temperature dependence of low field magnetic susceptibility was monitored with a CS-2 furnace apparatus attached to the KLY-3 susceptometer (Agico) from the *Laboratoire Géosciences Environnement Toulouse, Université de Toulouse III - Paul Sabatier*, using samples obtained from 19 rock specimens (Table 1). Fifteen samples were selected in order to represent all LOP granites and their different magnetic behavior according to previous magnetic susceptibility studies. Regarding the other post-tectonic plutons, 4 samples were selected (Table 1).

The samples were exposed to increasing temperatures up to 700°C in heating/ cooling cycles. The measurements were made in an argon flux in order to minimize oxygen fugacity, which reduces the oxidation of the samples and the consequent formation of new minerals with the increase and/or decrease of temperature. However, a small amount of oxygen is always present in the sample holder, so although the probability is lower, some mineralogical changes may still occur.

RESULTS AND DISCUSSION

Microscopy

The LOP granites are mostly composed of quartz, plagioclase (oligoclase, albite, and oligoclase-andesine), K-feldspars (orthoclase and microcline) and biotite (Table 2; Fig. 3A-C). As accessory minerals, muscovite I and/or II, zircon (Fig. 3A), sphene, allanite (Fig. 3D), fluorite (Fig. 3E), hematite \pm magnetite \pm ilmenite, chlorite \pm rutile, apatite, goethite, epidote, and tourmaline are present. In the BA granite, the muscovite is more abundant than in the LO and AC granites but is mostly of secondary origin resulting from deuteritic processes.

The plagioclases occur frequently zoned in all granites (Fig. 3C) and are albite-oligoclase in the BA granite, and oligoclase-andesine in the AC and LO granites. Occasionally, the plagioclases exhibit myrmekitic intergrowths with quartz, common mostly in the LO granite (Fig. 3F), and are altered by sericitization. The K-feldspar, perthitic orthoclase and microcline occur as heterogranular crystals. It is sometimes possible to observe some K-feldspar megacrystals with small plagioclase inclusions. The quartz has sometimes undulatory extinction, subgrain boundaries and fluid inclusion planes (Fig. 3G).

The biotite –brownish to greenish (Fig. 3E) in all granites– sometimes appears partially to completely chloritized. Occasionally, biotite is altered to muscovite, mainly in the leucocratic granite from BA, where biotite is observed in minor amounts (Fig. 3H). Frequently, biotite has pleochroic halos associated with zircon inclusions (Fig. 3A) and curved biotites are rare (Fig. 3H).

For a detailed magnetic mineralogy characterization, complementary studies of opaque minerals were conducted to point out the presence of magnetite, hematite and/or ilmenite. The SEM/EDS analyses showed the presence of different Fe-Ox minerals (Fig. 4) and other minerals in small amounts, like monazite, xenotime (Fig. 4B) and columbo-tantalite. Figure 4E and F are examples of some EDS spectra performed in LOP thin sections, showing the presence of Fe-Ox and Fe-Ox-Ti minerals. Figure 5 shows the micro-Raman spectrum of the principal magnetic minerals present in the LOP, namely magnetite (Fig. 5A), hematite (Fig. 5B) and ilmenite (Fig. 5C).

A summary of the petrographic descriptions of other Variscan granites made in previous studies is presented in Table 2. The Vila Pouca de Aguiar Pluton is located to the NE of the LOP, in the same NNE-SSW alignment, parallel to the Verin-Régua-Penacova Variscan fault. This pluton has two main granite facies with similar mineralogical composition. Both granites are composed

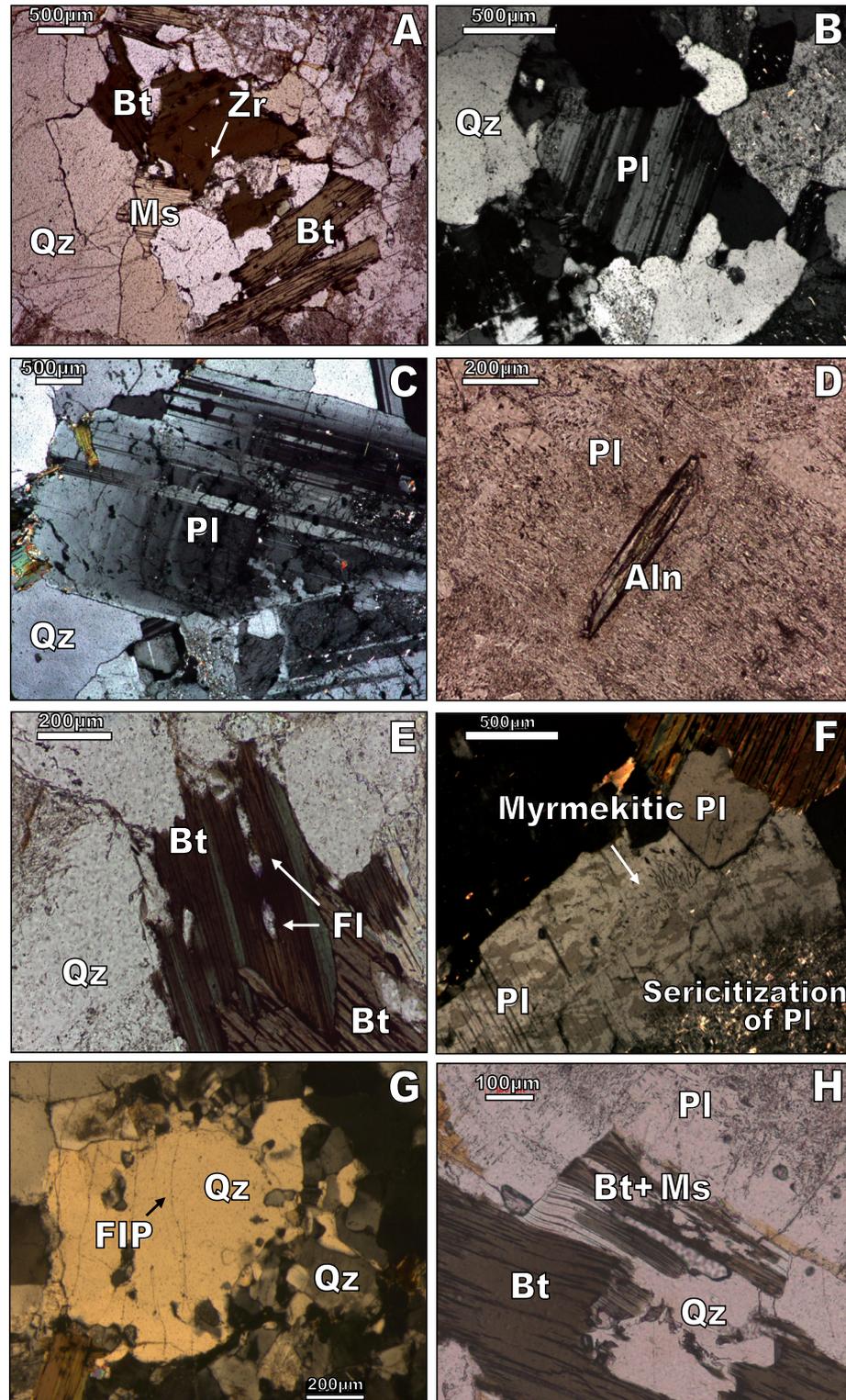


FIGURE 3. Photomicrographs of samples from the Lamas de Olo Pluton: A) Biotite and muscovite in plane-polarised light (Lamas de Olo granite); B, C) Plagioclase in crossed-polars (Barragem and Lamas de Olo granites, respectively); D) Alanite crystal on plagioclase in plane-polarised (Lamas de Olo granite); E) Biotite with brown and green colors cross-cutted by fluorite in plane-polarized (Lamas de Olo granite); F) Myrmekite texture and sericitization of plagioclase, in crossed-polars (Lamas de Olo granite); G) Microstructures in quartz, in crossed-polars (Lamas de Olo granite); H) Curved micas in plane-polarized (Lamas de Olo granite). Bt: biotite, Zr: zircon, Ms: muscovite, Qz: quartz, Pl: plagioclase, Aln: alanite, Fl: fluorite, FIP: fluid inclusion planes.

TABLE 2. Mineralogy and texture description summary for the studied granites. LOP: Lamas de Olo Pluton; VPAP: Vila Pouca de Aguiar Pluton; PGP: Peneda-Gerês Pluton; LMP: Lavadores-Madalena Pluton. According to this study data and previous studies carried out by: Sant’Ovaia *et al.* (2000) and Martins and Noronha (2006) for VPAP; Mendes and Dias (2004) and Cottard (1979) for PGP; Martins *et al.* (2011) and Sant’Ovaia *et al.* (2014) for LMP

Pluton	$^{207}\text{Pb}/^{235}\text{U}$ dating	Granite	Texture	Mineralogy
LOP	~297Ma (Fernandes <i>et al.</i> , 2013)	Lamas de Olo	Medium to coarse-grained porphyritic biotite granite	Quartz, plagioclase, K-feldspar and biotite; zircon, sphene, allanite, fluorite, hematite, magnetite, ilmenite, chlorite, rutile, apatite, goethite, epidote, tourmaline, xenotime, monazite and rare muscovite I and/or II
		Alto dos Cabeços	Fine to medium-grained porphyritic biotite granite	Quartz, plagioclase, K-feldspar and biotite; muscovite I and/or II, zircon, sphene, allanite, fluorite, hematite, ilmenite, chlorite, rutile, apatite, goethite, epidote, tourmaline, xenotime and monazite
		Barragem	Leucocratic fine to medium-grained and slightly porphyritic biotite granite	Quartz, perthitic K-feldspar, biotite and plagioclase; zircon, apatite, allanite, xenotime, ilmenite, sphene and rare monazite
VPAP	~297-298Ma (Martins <i>et al.</i> , 2009)	Pedras Salgadas	Porphyritic fine to medium-grained biotite monzogranite	Quartz, perthitic K-feldspar, plagioclase, quartz and biotite; ilmenite, zircon, apatite, allanite, sphene, monazite and scarce hornblende
		Vila Pouca de Aguiar	Porphyritic medium to coarse-grained biotite monzogranite	Quartz, perthitic K-feldspar, biotite and plagioclase; zircon, apatite, allanite, xenotime, ilmenite and sphene
PGP	~290-296Ma (Dias <i>et al.</i> , 1998)	Gerês	Porphyritic medium to coarse-grained biotite granite	Perthitic K-feldspar, plagioclase, quartz and biotite; ilmenite, zircon, apatite, allanite, sphene, monazite and scarce hornblende
		Illa	Fine-grained two-mica monzogranite	Perthitic K-feldspar, plagioclase, quartz and biotite; muscovite I and/or II, ilmenite, zircon, apatite and monazite
		Carris	Medium to coarse-grained biotite granite	Perthitic K-feldspar, plagioclase, quartz and biotite; magnetite, ilmenite, zircon, apatite and monazite
LMP	~298Ma (Martins <i>et al.</i> , 2011)	Madalena	Porphyritic medium to coarse-grained biotite granite with porphyritic texture	Quartz, K-feldspar, plagioclase and biotite; magnetite, ilmenite, hematite, zircon, apatite, muscovite and chlorite
		Lavadores	Porphyritic coarse-grained biotite granite	Quartz, plagioclase, perthitic K-feldspar and biotite; magnetite, ilmenite, hematite, zircon, sphene, apatite, allanite and amphibole

of quartz, perthitic K-feldspar (orthoclase and microcline) and plagioclase with normal zoning. The Vila Pouca de Aguiar granite contains also oligoclase-andesine and the Pedras Salgadas granite albite-oligoclase. The biotite, the only ferromagnesian phase, is more abundant in the Vila Pouca de Aguiar granite. Accessory minerals include zircon, apatite, allanite, xenotime, ilmenite, sphene and rare monazite (Martins and Noronha, 2006; Sant’Ovaia *et al.*, 2000).

The three granites of the Peneda-Gerês Pluton, the northern pluton, are similar in their mineralogy. They contain perthitic K-feldspar, sometimes as megacrystals, plagioclase (albite-oligoclase, rarely andesine), quartz, biotite, ilmenite, zircon, apatite, and monazite. Some minerals are specific to some granites: allanite, sphene, and hornblende to the Gerês granite; muscovite to the Illa granite, resulting from later muscovitization; and magnetite to the Carris granite (Cottard, 1979; Mendes and Dias, 2004).

The Lavadores-Madalena Pluton is composed of two similar porphyritic, coarse-grained biotite granites. The Lavadores granite consists of quartz, plagioclase (andesine and oligoclase), perthitic K-feldspar (orthoclase and microcline), biotite, magnetite, ilmenite, hematite, zircon, sphene, apatite, allanite, and amphibole. The Madalena granite is composed of quartz, K-feldspar (orthoclase), plagioclase and biotite. Zircon, apatite, muscovite, chlorite, magnetite, ilmenite and hematite are present as accessory minerals (Martins *et al.*, 2011; Sant’Ovaia *et al.*, 2014).

Magnetic mineralogy

Magnetic susceptibility and isothermal remanent magnetization data

The magnetic susceptibility study in the post-D₃ Variscan granites showed the presence of two types of magnetic behavior: ilmenite- and magnetite-type, with the predominance of the ilmenite-type ($K_m < 1,000 \mu\text{SI}$) (Table 3; Fig. 6). Previous studies in the LOP (Cruz *et al.*, 2016) indicated that: i) the K_m values in the LO granite have a huge variability (ranging from low to high K_m), ii) the AC granite has intermediate K_m values, and iii) the BA granite has the lower K_m values (Table 3). This data show that LO granite is the most heterogeneous granite of the pluton, suggesting that both ferromagnetic and paramagnetic behavior are present. On the other hand, the AC and BA granites have exclusively a paramagnetic behavior. So, the LOP is the most heterogeneous pluton, showing three K_m classes, below $50 \mu\text{SI}$, between $50 \mu\text{SI}$ and $2,000 \mu\text{SI}$, and higher than $2,000 \mu\text{SI}$ (Table 3; Fig. 6B). The Vila Pouca de Aguiar Pluton has minor K_m dispersion with values below $300 \mu\text{SI}$, showing a typical ilmenite-type behavior (Sant’Ovaia *et al.*, 2013a; Table 3; Fig. 6B). In the Peneda-Gerês Pluton, two granite facies are ilmenite-type and one is magnetite- and ilmenite-type. The granites from the Lavadores-Madalena Pluton have both minor K_m dispersion and the highest K_m mean values, always with K_m higher than $1,550 \mu\text{SI}$ (Martins *et al.*, 2011; Sant’Ovaia *et al.*, 2014; Table 3; Fig. 6B).

The IRM data from the LOP are presented in Table 4 and Figure 7. The LM 5 sample shows saturation at

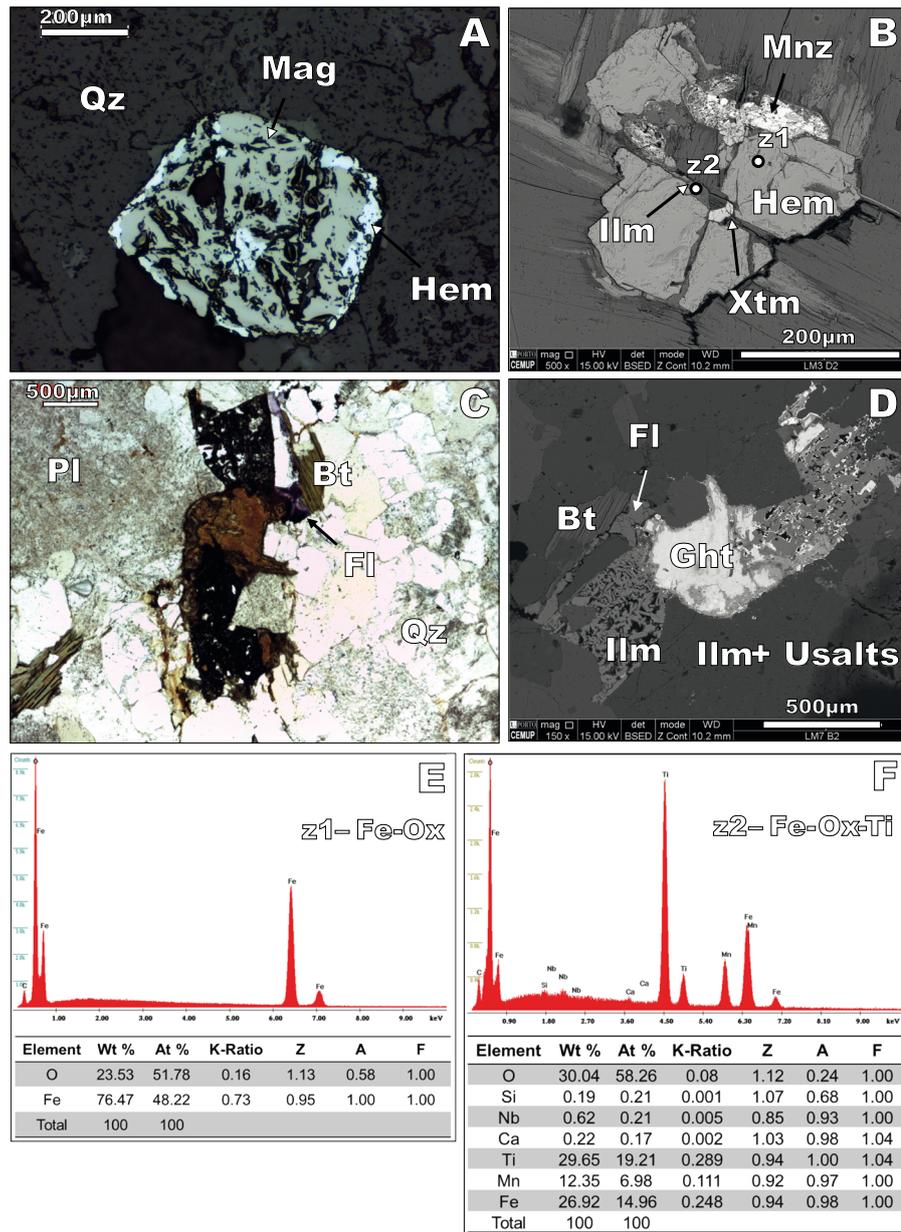


FIGURE 4. Photomicrographs (A-D) and EDS spectra of Lamas de Olo Pluton (E-F). A) Individual grain with magnetite and hematite, in reflected and plane-polarized light (Lamas de Olo granite); B) SEM image of Fe-Ox and other minerals in minor amounts with location of the EDS spectra z1 and z2 (Lamas de Olo granite); C) A mixture of Fe-Ox minerals in transmitted and plane-polarized light (Barragem granite); D) SEM image of the same textures shown in (C); E) EDS spectrum of a Fe-Ox mineral (z1) with semi-quantitative analysis; F) EDS spectrum of a Fe-Ox-Ti mineral (z2) with semi-quantitative analysis. Mag: magnetite, Hem: hematite, Qz: quartz, Bt: biotite, Mnz: monazite, Ilm: ilmenite, Xtm: xenotime, Pl: plagioclase, Fl: fluorite, Ght: goethite, U salts: uranium salts; Fe-Ox: iron oxide; Fe-Ox-Ti: iron oxide with titanium.

fields *ca.* 338mT, suggesting the dominance of low coercive minerals, while LM 16 does not reach saturation at the maximum imparted field. The $S_{0.3T}$ for these samples is 0.999 and 0.839, respectively, confirming the magnetic heterogeneity of the LO granite, with areas of predominance of low coercive minerals and areas where ferrimagnetic minerals are scarce. The curve for sample LM 26, from the AC granite, indicates higher amounts of ferrimagnetic minerals ($S_{0.3T} = 0.961$), however does

not show saturation at the applied field. The acquisition curve at *ca.* 950mT is apparently stable, however, a zoom shows a small increase in intensity with increasing fields, indicating a small amount of a high coercive mineral such as hematite. The LM 7 sample, from the BA granite, shows some oscillations in the intensity with increasing fields due to a weak magnetic signal. However, the $S_{0.3T}$ from this sample is 0.995, suggesting the presence of low coercive minerals.

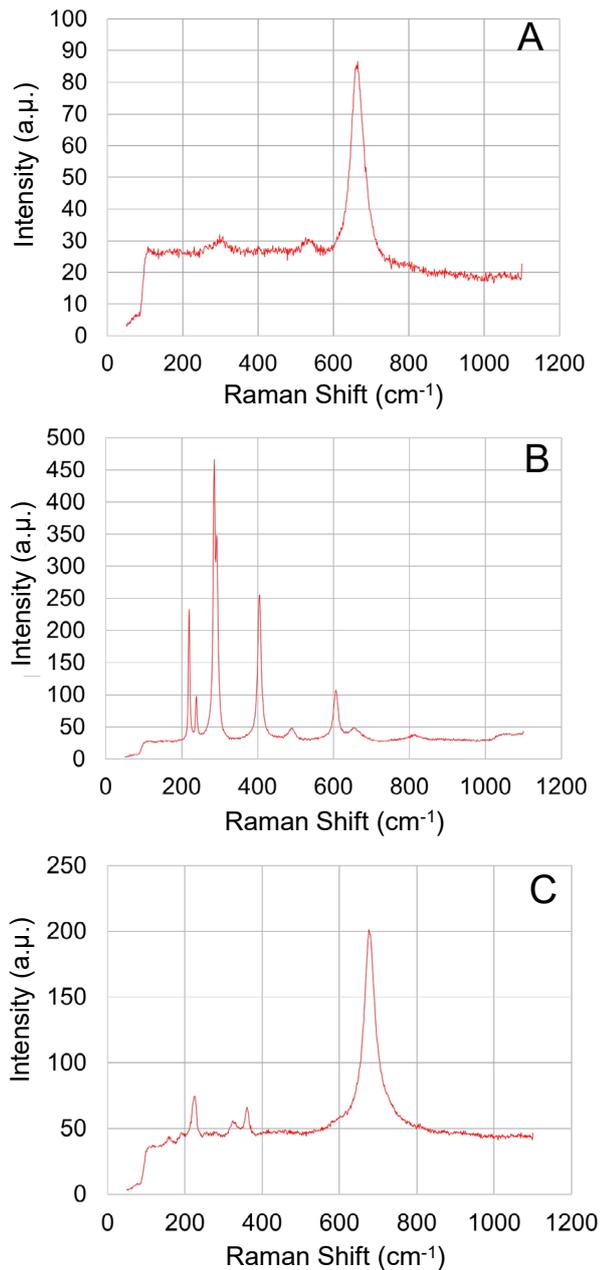


FIGURE 5. Micro-Raman spectra of Fe-Ox minerals from the Lamas de Olo Pluton: A) Magnetite; B) Hematite; C) Ilmenite.

After analysing the data using the CLG function (Kruiver *et al.*, 2001; Robertson and France, 1994), best fits of the raw IRM curves are obtained by considering two or three components (Table 4; Fig. 7). All specimens have, at least two components: component 1 with low coercive phase and $B_{1/2}$ (the field at which half of the SIRM is reached) *ca.* 17mT and component 2, showing an intermediated coercive phase and $B_{1/2}$ *ca.* 59mT. Two samples show a third component, with higher coercivity with $B_{1/2}$ *ca.* 466mT.

The Dispersion Parameter (DP) ranges between 0.25 and 0.42 for all components, with mean values of 0.36, 0.31 and 0.34 for component 1, 2 and 3, respectively. According to Dunlop and Özdemir (1997), the coercivity increases with the replacement of Fe^{3+} by Ti in titanomagnetite. Taking this into account, the components 1, 2 and 3 have been interpreted as magnetite, Ti-poor magnetite and hematite, respectively (Abrajevitch and Kodama, 2011; Font *et al.*, 2014; Kruiver *et al.*, 2001; Maxbauer *et al.*, 2016; Robertson and France, 1994). Ti-poor magnetite (component 2) is dominant in all samples (contributes to more than 50% of the total remanence), with exception of LM 7 with 13% of Ti-poor magnetite and 87% of magnetite (component 1; Table 4).

According to some authors (*e.g.* Font *et al.*, 2009; Font *et al.*, 2014; Kruiver *et al.*, 2001), the differences in $B_{1/2}$ and DP values can be interpreted as: varying degrees of oxidation of magnetite for component 1; different Ti-content in the Ti-poor magnetite for component 2; and different magnetic grain sizes for component 3.

IRM studies carried out by Sant’Ovaia (1993) in the Vila Pouca de Aguiar Pluton showed that the Vila Pouca de Aguiar and Pedras Salgadas granites have similar magnetic behavior. The $S_{0.3T}$ ranges between 0.624 and 0.968, with minor values obtained in the Pedras Salgadas granite. Both granites show an absence of saturation, however, the Pedras Salgadas granite presents an acquisition curve with a high slope, showing the depletion of ilmenite (Table 4; Fig. 8). Previous studies in the Lavadores-Madalena Pluton showed saturation fields under 300mT and $S_{0.3T}$ values higher than 0.955 (Table 4; Fig. 8), pointing out high ferromagnetic mineral content (Martins *et al.*, 2011; Sant’Ovaia *et al.*, 2014).

Frequency-dependent susceptibility (KfD%) studies were made in LOP samples. In the BA granite, KfD% was higher than 10% (LM 7 ~ 35%; LM 58 ~ 29% and LM 62 ~ 30%), which have been interpreted by some authors (*e.g.* Dearing *et al.*, 1996) as due to the presence of superparamagnetic fine grains, a weak magnetic signal or a high alteration degree. On the other hand, the values acquired for the LO samples (KfD% < 2%) indicated absence of superparamagnetic grains, and for the AC samples (4% < KfD% < 7%) an admixture of superparamagnetic and coarser non-superparamagnetic grains (Dearing *et al.*, 1996). The presence of superparamagnetic minerals in samples from the BA granite could explain the lower values of the magnetic susceptibility (mean K_m *ca.* 27 μ SI) and the presence of low coercive minerals.

Thermomagnetic experiments

Figures 9 and 10 show different magnetic susceptibility-temperature curves (Kt vs T charts) obtained from samples

TABLE 3. Magnetic mineralogy of studied granites according to this study data and previous studies carried out by: Cruz et al. (2016) for LOP; Sant’Ovaia (1993) and Sant’Ovaia et al. (2013) for VPAP; Sant’Ovaia et al. (2013) for PGP; Martins et al. (2011) and Sant’Ovaia et al. (2014) for LMP. The maghemite appears between brackets because it is a neoformation mineral due to oxidation of magnetite during thermomagnetic experiments. LOP: Lamas de Olo Pluton; VPAP: Vila Pouca de Aguiar Pluton; PGP: Peneda-Gerês Pluton; LMP: Lavadores-Madalena Pluton, and associated mineralizations; N: number of sampling sites; n: number of samples; n.a.: not attributed

Pluton	Granite	Magnetic behavior	K_m (μ SI)				Magnetic mineralogy		Associated mineralizations
			mean	range	N	n	Magnetic susceptibility studies and thermomagnetic experiments	Cumulative log-Gaussian analysis	
LOP	Lamas de Olo	Magnetite- and ilmenite-type	2,078	67-44,382	39	293	Ti-poor magnetite + hematite + pure magnetite + ilmenite (+ maghemite)	Ti-poor magnetite > magnetite > hematite	W-Mo
	Alto dos Cabeços	Ilmenite-type	334	125-804	4	29	Ti-poor magnetite + ilmenite + hematite + magnetite (+ maghemite)	Ti-poor magnetite > magnetite > hematite	
	Barragem	Ilmenite-type	27	21-34	5	28	Hematite + ilmenite + Ti-poor magnetite + magnetite	Magnetite > Ti-poor magnetite	
VPAP	Pedras Salgadas	Ilmenite-type	68	44-121	35	280	Hematite + ilmenite + Ti-poor magnetite	n.a	Without mineralizations
	Vila Pouca de Aguiar	Ilmenite-type	135	60-218	81	648	Hematite + ilmenite + Ti-poor magnetite	n.a	
PGP	Gerês	Magnetite- and ilmenite-type	463	46-1,000	7	7	Hematite + Ti-poor magnetite + ilmenite	n.a	W-Mo Mo W-Sn-Mo-Bi
	Illa	Ilmenite-type	81	55-109	3	3	Hematite + ilmenite + Ti-poor magnetite	n.a	
	Carris	Ilmenite-type	33	29-36	3	3	Hematite + Ti-poor magnetite + ilmenite	n.a	
LMP	Madalena	Magnetite-type	7,037	1,550-9,920	7	80	Ti-poor magnetite + ilmenite + hematite (+ maghemite)	n.a	W-Mo
	Lavadores	Magnetite-type	15,971	10,589-19,303	6	40	Magnetite + ilmenite + hematite (+ maghemite)	n.a	

of the Variscan granites. It should be noted that in the K_t vs T charts the susceptibility data, on the y-axis, is referred to as total or bulk susceptibility (K_t), and it is not corrected by volume or mass. These values are, therefore, quite different from the K_m values measured in the magnetic susceptibility experiments, where the K_m values are corrected for the specimen volume or mass.

The K_t vs T curves from selected LO granite samples show a significant fall of K_t at ca. 580°C, indicating magnetite T_C (Fig. 9A). Other curves have a K_t drop just before 580°C corresponding to a Ti-poor magnetite (Fig. 9B, C, D). Some samples have a slighter drop than the others

because they have a smaller Ti-poor magnetite content (Fig. 9D). The K_t vs T curves from BA samples show the presence of hematite (Fig. 9E, F, G), and some traces of Ti-poor magnetite. The K_t values in the AC granite are, in general, lower than in the LO granite, but some K_t vs T curves show a slight drop just before 580°C, pointing to an ilmenite-type granite, but probably with some traces of Ti-poor magnetite (Fig. 9H, I). This could be related to the circulation of post-magmatic fluids capable of changing the characteristics of the primary magnetic minerals with the formation of new, less magnetic minerals, like hematite, maghemite or goethite (Lagoeiro, 1998; Nédélec et al., 2015). All the LOP curves show the presence of hematite and, in some

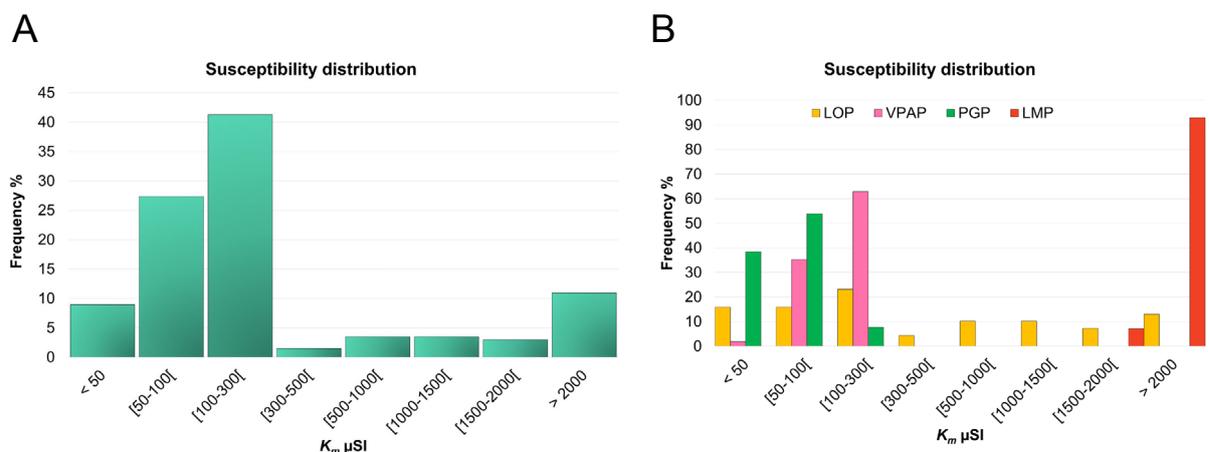


FIGURE 6. Relative frequency magnetic susceptibility histograms: A) for all the studied granites; B) for the plutons; LOP: Lamas de Olo Pluton; VPAP: Vila Pouca de Aguiar Pluton, PGP: Peneda-Gerês Pluton, LMP: Lavadores-Madalena Pluton.

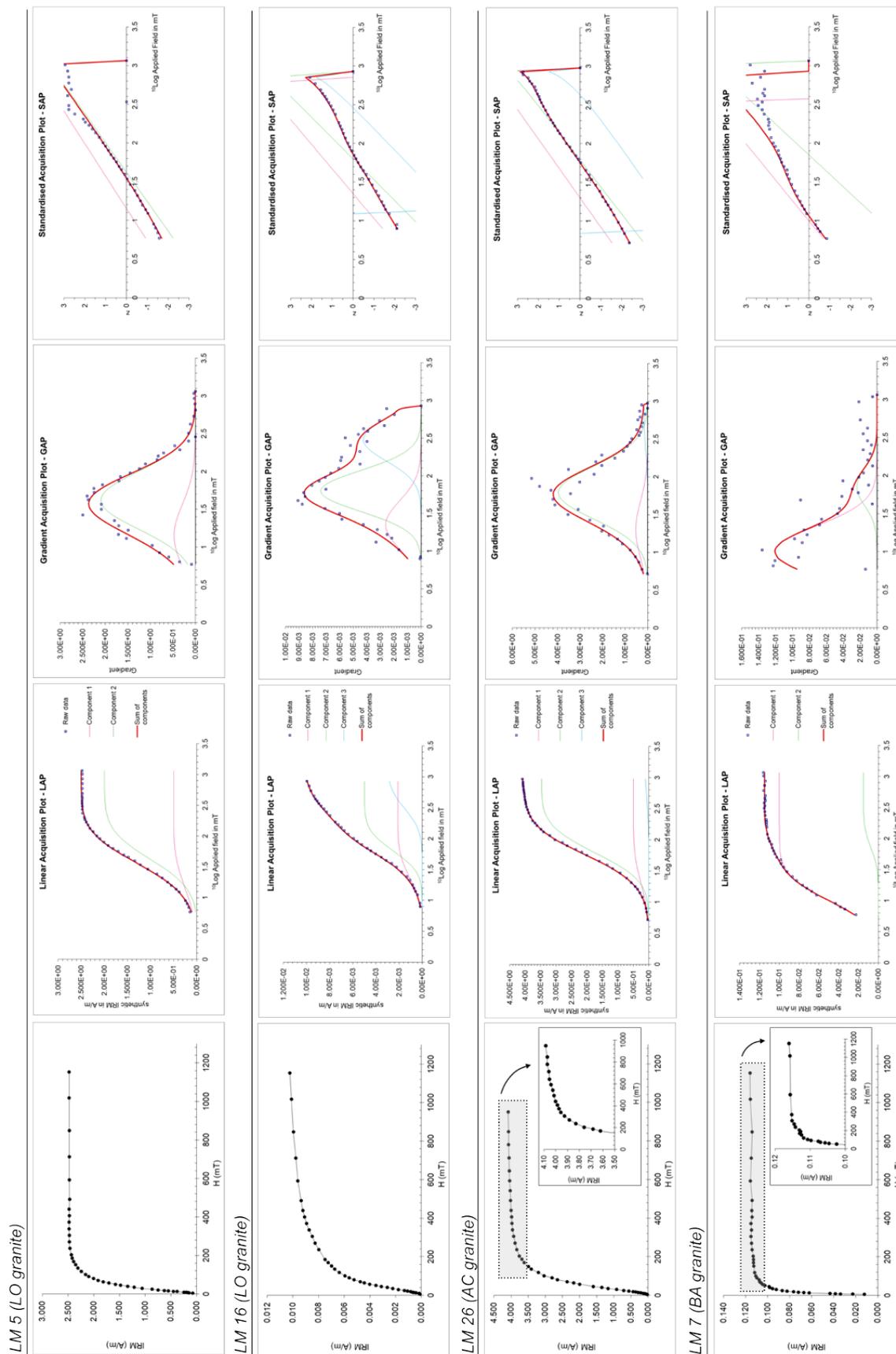


FIGURE 7. IRM acquisition curves for representative samples from the Lamas de Olo (LO) granite; LM 26 from the Alto dos Cabeços (AC) granite (with zoom showing the absence of saturation); LM 7 from the Barragem (BA) granite (with zoom showing the absence of saturation). Data were analyzed using a Cumulative Log-Gaussian (CLG) function using the software developed by [Kruiver et al. \(2001\)](#).

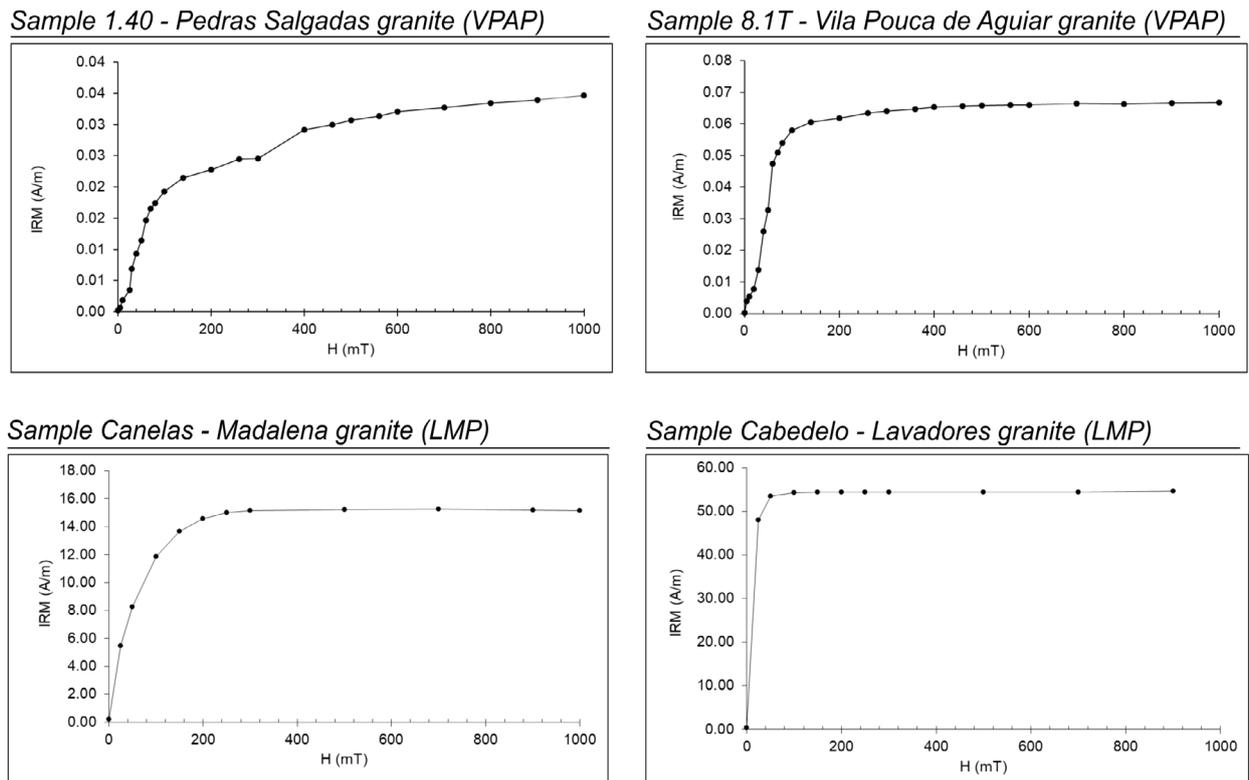


FIGURE 8. Examples of IRM acquisition curves from the Vila Pouca de Aguiar Pluton (VPAP) and the Lavadores-Madalena Pluton (LMP) from earlier studies (Martins et al., 2011; Sant’Ovaia et al., 1993; Sant’Ovaia et al., 2014).

TABLE 4. Isothermal remanent magnetization data from the studied plutons. Vila Pouca de Aguiar Pluton (VPAP) and Lavadores-Madalena Pluton (LMP) data previously published in Sant’Ovaia (1993) and in Martins et al. (2011) and Sant’Ovaia et al. (2014), respectively. IRM_{1T}: Isothermal remanent magnetization at 1T; IRM_{0.3T}: isothermal remanent magnetization at 300mT. The Cumulative log-Gaussian (CLG) analysis were performed only in the LOP samples; CLG treatment allows defining for each magnetic component, %: percent contribution; SIRM: values of saturation of isothermal remanent magnetization; B_{1/2}: the field at which half of the SIRM is reached; DP: dispersion parameter; S_{0.3T} = IRM_{0.3T}/IRM_{1T} (Robertson and France, 1994). The CLG treatment was made using the software developed by Kruiver et al. (2001). n.a.: not attributed

Pluton	Granite	Sample	IRM _{1T} (A/m)	IRM _{0.3T} (A/m)	S _{0.3T}	CLG analysis											
						Component 1 (20-40 mT)				Component 2 (50-80 mT)				Component 3 (100-250 mT)			
						%	SIRM (A/m)	B _{1/2} (mT)	DP	%	SIRM (A/m)	B _{1/2} (mT)	DP	%	SIRM (A/m)	B _{1/2} (mT)	DP
LOP	Lamas de Olo	LM 5	2.480	2.477	0.999	20	0.50	14.13	0.42	80	2.00	41.69	0.38				
		LM 16	0.010	0.008	0.839	21	0.00	22.39	0.32	50	0.01	63.10	0.27	30	0.00	302.00	0.28
	Alto dos Cabeços	LM 26	4.086	3.925	0.961	12	0.50	19.95	0.38	84	3.47	61.66	0.35	4	0.17	630.96	0.4
		Barragem	LM 7	0.116	0.115	0.995	87	0.10	10.00	0.33	13	0.02	70.79	0.25			
VPAP	Pedras Salgadas	1.40	0.035	0.025	0.706												
		4.2T	0.015	0.009	0.624												
		13.4T	0.024	0.022	0.883												
	Vila Pouca de Aguiar	12.20	0.021	0.019	0.881												
		8.1T	0.067	0.064	0.959												
		11.4T	0.061	0.059	0.968					n.a.							
LMP	Madalena	Cavadão	13.313	13.273	0.997												
		Canelas	15.160	15.158	1.000												
		Carvalhos	53.520	52.786	0.986												
		Saúde	16.618	15.870	0.955												
		Pedroso	32.060	33.566	1.047												
		Lavadores	Cabedelo	54.621	54.435	0.997											

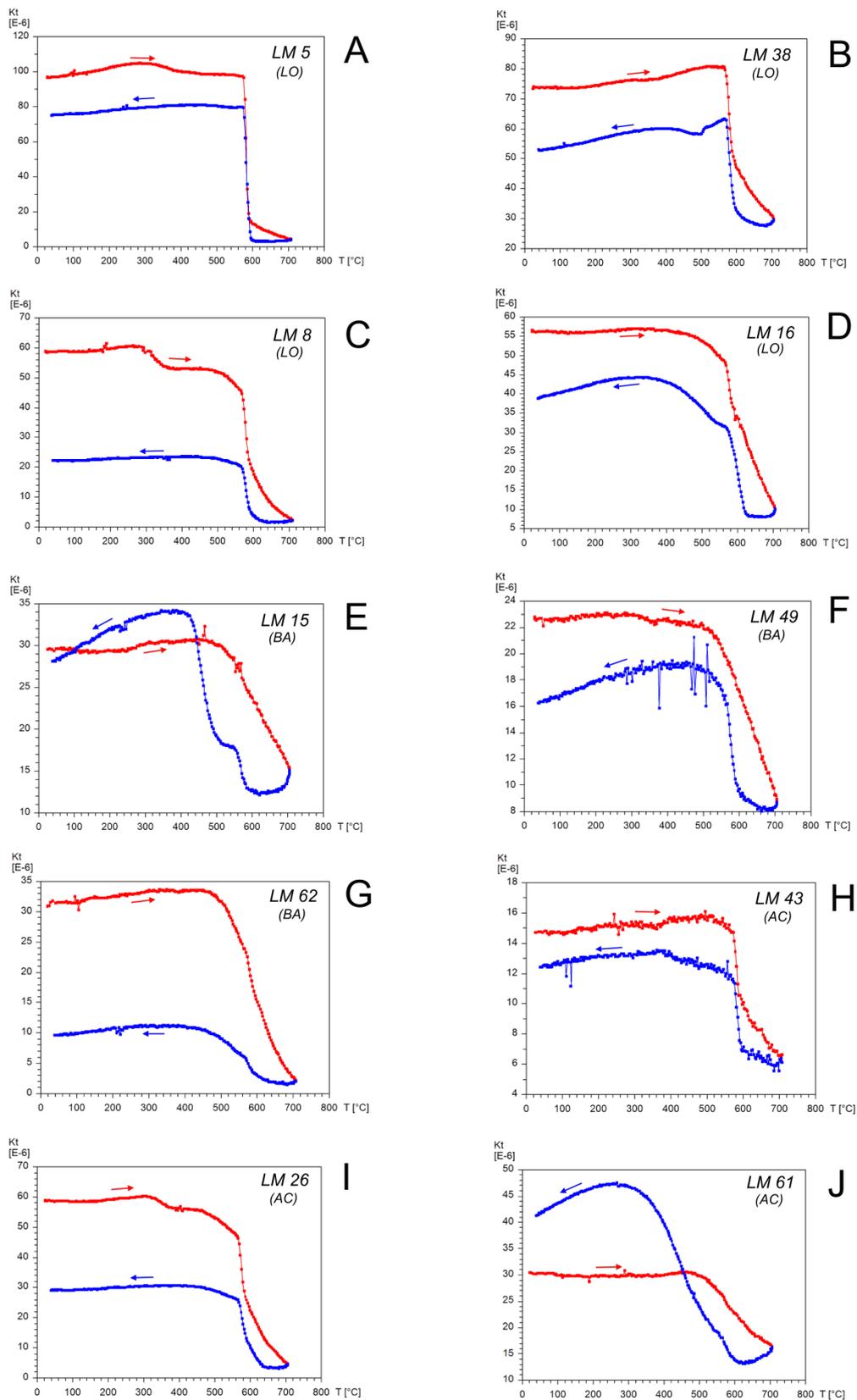


FIGURE 9. Normalized K_t vs T during heating (red arrow toward right) and cooling (blue arrow toward left) from representative samples of LOP granites: A-D) from the Lamas de Olo (LO) granite; E-G) from the Barragem (BA) granite; H-J) from the Alto dos Cabeços (AC) granite. Note that the y-axes are not on the same scale in all charts, so it is necessary to consider the K_t values themselves.

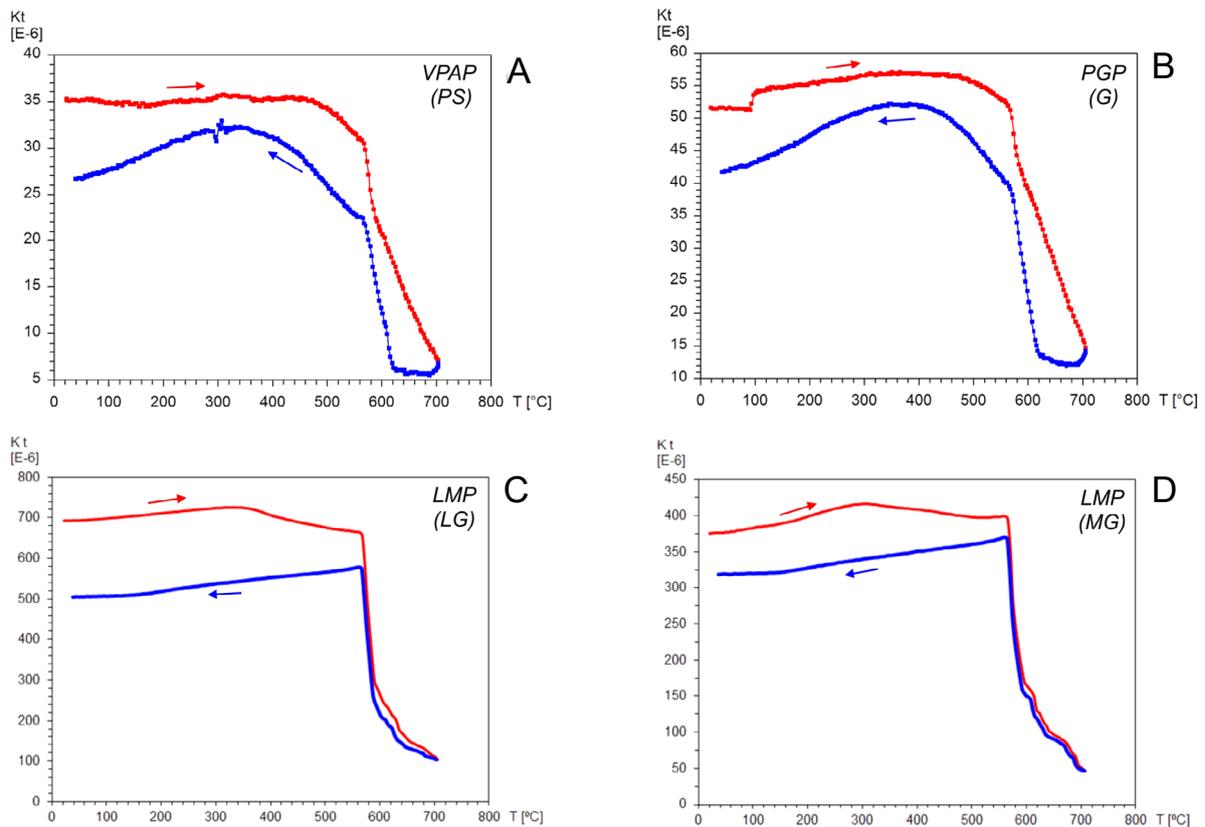


FIGURE 10. Normalized K_t vs T during heating (red arrow toward right) and cooling (blue arrow toward left) from Variscan granites other than the LOP granites: A) Pedras Salgadas granite from the Vila Pouca de Aguiar Pluton (VPAP-PS); B) Gerês granite from the Peneda-Gerês Pluton (PGP-G); C) Lavadores granite from the Lavadores-Madalena Pluton (LMP-LG) and D) Madalena granite from the Lavadores-Madalena Pluton (LMP-MG).

cases (Fig. 9E, J), during the cooling cycle, below 580°C, K_t progressively increases, surpassing the values attained in the heating curve, suggesting the formation of a magnetic mineral, probably Ti-poor magnetite.

In some K_t vs T curves from the LO and AC granites, an inflection around 300°C was observed on the heating curves (Fig. 9A, C, D). This inflection may correspond to pyrrhotite or to maghemite. The pyrrhotite is a ferromagnetic mineral, with T_C at ca. 320°C (Butler, 1992), and is very common in granites (e.g. Jover *et al.*, 1989); however, the monoclinic pyrrhotite is chemically stable up to its Curie point (Dekkers, 1989), which cannot explain the disappearance of the inflection in the cooling cycling. The presence of maghemite in the analyzed samples, may be due to the oxidation of magnetite with increased furnace temperatures and the presence of oxygen in the system (Evans and Heller, 2003). Maghemite is normally destroyed before 350°C, which could explain the inflection around 300°C in the heating curves and its absence in the cooling cycle. For that reason, maghemite is probably the magnetic mineral responsible for the inflection at 320°C.

The thermomagnetic behavior of other Variscan granites is quite similar to the LOP. With the exception

of the Lavadores-Madalena Pluton, the thermomagnetic curves show a slight drop before 580°C, indicating a small quantity of magnetic minerals. The K_t drops around 570°C and 565°C in the Vila Pouca de Aguiar Pluton (Fig. 10A) and Peneda-Gerês Pluton (Fig. 10B) samples, respectively. These values indicate the presence of Ti-poor magnetite, although Ti content in the sample from Peneda-Gerês Pluton is slightly higher (lower T_C) than the Vila Pouca de Aguiar Pluton. Regarding the Lavadores-Madalena Pluton experiments, the well-marked drop at 580°C demonstrates the presence of magnetite and the inflection at 300°C in the heating curve suggests the presence of maghemite in the Lavadores (Fig. 10C) and Madalena (Fig. 10D) granites. All K_t vs T curves also revealed the presence of hematite.

FINAL REMARKS AND CONCLUSIONS

The magnetic mineralogy studies in the Portuguese post-tectonic Variscan granites allowed to conclude that these are ilmenite- and magnetite-type-granites. Thermomagnetic curves are quite similar and show the presence of magnetite/Ti-poor magnetite and hematite in all the plutons, though in different proportions. When the magnetite content is compared, a sequence can be established: Lavadores-

Madalena>Lamas de Olo>Peneda-Gerês>Vila Pouca de Aguiar. The presence of magnetite as accessory mineral allowed inferring the generation depth of magmas.

Previous studies in the LOP granites, demonstrated the co-occurrence of ferromagnetic and paramagnetic behavior in the pluton (Cruz *et al.*, 2016). Alteration associated with the circulation of post-magmatic fluids, could explain the presence of other oxides such as Ti-poor magnetite and hematite in the granites.

The complexity of the magnetic mineralogy is not recognized in the magnetic susceptibility measurements, because this parameter only expresses the combined magnetic behavior of all minerals. For that reason, other methodologies (*e.g.* IRM, thermomagnetic curves, and frequency-dependent susceptibility) should be performed to identify magnetic minerals and their complex magnetic behaviors.

In the Vila Pouca de Aguiar Pluton, the magnetic mineralogy is simpler when compared with other plutons. Studies carried out in the Peneda-Gerês and Lavadores-Madalena plutons pointed to a complex mineralogy, with the presence of both magnetite and ilmenite mineralogy in their granites.

With the exception of the Vila Pouca de Aguiar Pluton (where mineralizations have not yet been observed), the Variscan granites have associated mineralizations. The LOP has W (Mo) mineralizations associated with the AC and LO granites (Helal, 1992; Pereira, 1989); the Peneda-Gerês Pluton has W (Sn-Mo-Bi) mineralizations associated with their granites (Cottard, 1979; Cheilletz and Giuliani, 1982; Noronha, 1984), and in the Lavadores-Madalena Pluton, W (Mo) local occurrences have been described associated with the Madalena granite (Teixeira and Perdigão, 1962).

The Lamas de Olo, Peneda-Gerês and Lavadores-Madalena plutons are all post-tectonic and composed of several granites of complex magnetic mineralogy, with magnetite/Ti-poor magnetite and ilmenite. Then we conclude that the presence of ilmenite- and magnetite-type in the same pluton, results from complex redox reactions, that occurs during the ascend and granite emplacement. The presence of magnetite pointing to a deep magma origin and oxidizing conditions in the magma chamber. This is an important metallogenic indicator of the W (Mo) mineralizations associated with post-tectonic biotite granites. Thus, the magnetic mineralogy can be a useful pathfinder for the W (Mo) mineralization exploration in the Iberian Peninsula.

The measurement of magnetic susceptibility is an unexpensive and rapid technique that can be used to

characterize the magnetic behavior. The presence of magnetite and/or ilmenite allows to infer the redox condition of the granitic magma and to assess the metallogenic potential. However, the low field magnetic susceptibility measurements have some limitations when complex magnetic mineralogy is present, and for that reason, other magnetic studies should be performed.

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